

# NEAR-REAL-TIME MONITORING OF GLOBAL OCEAN CARBON SINK

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## ABSTRACT

The ocean, absorbing about 25% of anthropogenic  $CO_2$  emissions, plays a crucial role in mitigating climate change. However, the delayed (by one year) traditional estimates of ocean-atmosphere  $CO_2$  flux hinder timely understanding and response to the global carbon cycle's dynamics. Addressing this challenge, we introduce Carbon Monitor Ocean (CMO-NRT), a pioneering dataset providing near-real-time, monthly gridded estimates of global surface ocean fugacity of  $CO_2$  ( $fCO_2$ ) and ocean-atmosphere  $CO_2$  flux from January 2022 to July 2023. This dataset marks a significant advancement by updating the global carbon budget's estimates through a fusion of data from 10 Global Ocean Biogeochemical Models (GOBMs) and 8 data products into a near-real-time analysis framework. By harnessing the power of Convolutional Neural Networks (CNNs) and semi-supervised learning techniques, we decode the complex nonlinear relationships between model or product estimates and observed environmental predictors. The predictive models, both for GOBM and data products, exhibit exceptional accuracy, with root mean square errors (RMSEs) maintaining below the 5% threshold. This advancement supports more effective climate change mitigation efforts by providing scientists and policymakers with timely and accurate data.

## 1 INTRODUCTION

The ocean is a pivotal component in the Earth's climate system, acting as a major sink for anthropogenic heat and carbon dioxide (CO<sub>2</sub>). This crucial role underscores the necessity for timely and accurate estimates of the global ocean carbon sink to inform climate change mitigation efforts and support the global stocktake process under the Paris Climate Agreement. Traditionally, the annual

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Global Carbon Budget report has provided estimates of the global ocean carbon sink, yet these figures are historically delayed by approximately one year Friedlingstein et al. (2022) due to computational and data gathering constraints. This latency hinders the timely assessment and response to the changing state of global carbon sinks, emphasizing the need for more immediate data solutions.

In response to this need, numerous methodologies have been developed, utilizing both in situ measurements and global ocean biogeochemical models (GOBMs) to estimate surface ocean fugacity of  $CO_2$  ( $fCO_2$ ), surface ocean partial pressure of  $CO_2$  ( $pCO_2$ ) and air-sea  $CO_2$  flux. The Surface Ocean  $CO_2$  Atlas ([www.socat.info](http://www.socat.info)) (SOCAT), a community-led database, has played a central role by providing a comprehensive repository of quality-controlled surface ocean  $fCO_2$  measurements. These data, spanning from 1957 to 2022 in its latest update Bakker et al. (2016), form the backbone of observation-based products that estimate  $pCO_2$  across the globe. These products leverage sparse  $pCO_2$  observations from SOCAT, applying multivariate linear regression or machine learning algorithms alongside observations of related variables to estimate  $pCO_2$  at any given location and times (Landschützer et al., 2016; Rödenbeck et al., 2022; Chau et al., 2022; Gloege et al., 2022; Watson et al., 2020; Zeng et al., 2014; Iida et al., 2021; Gregor & Gruber, 2021). Parallelly, GOBMs offer a holistic simulation of the ocean's carbonate system by integrating its physical, biological, and chemical dynamics (Wright et al., 2021; Schwinger et al., 2016; Lacroix et al., 2021; Berthet et al., 2019; Hauck et al., 2020; Liao et al., 2020; Doney et al., 2009; Aumont et al., 2015; Nakano et al., 2011; Urakawa et al., 2020; Long et al., 2021). Though these models deliver extensive insights into the ocean's carbon cycle, their utility is constrained by computational demands and the inherent delay in SOCAT data updates, culminating in a significant latency in current global ocean carbon sink estimates.

To bridge this gap, we introduce Carbon Monitor Ocean (CMO-NRT), a dataset providing near-real-time, monthly gridded global surface ocean  $fCO_2$  and ocean-atmosphere  $CO_2$  flux data from January 2022 to July 2023. CMO-NRT represents a paradigm shift in the monitoring of oceanic carbon by employing a deep learning approach that amalgamates temporal, spatial, and environmental variables, offering a timely alternative to the delayed updates characteristic of existing methods. This dataset not only addresses the critical need for near-real-time data but also showcases the potential of advanced computational techniques to enhance our understanding of global biogeochemical cycles. This paper details the development and validation of CMO-NRT, illustrating its methodology, data integration processes, and the implications of its findings for global carbon monitoring efforts.

## 2 DATA AND METHODS

Carbon Monitor Ocean (CMO-NRT) utilizes advanced deep learning methods to build models that connects the output of GOBMs or ocean data products with relevant environmental variable observations. We then use CMO-NRT to extend the monthly gridded global surface ocean  $fCO_2$  and ocean-atmosphere  $CO_2$  flux estimates from each of the 10 GOBMs and 8 data products used in the Global Carbon Budget 2022, from January 2022 to July 2023. A pipeline of CMO-NRT is shown in Figure 1. Below we describe the calculation process in detail.

### 2.1 DATA SOURCES AND PRE-PROCESSING

#### 2.1.1 GLOBAL OCEAN BIOGEOCHEMICAL MODELS AND OBSERVATION-BASED DATA PRODUCTS

We utilize monthly data until the end of 2021 from ten GOBMs and eight data products contributing to the Global Carbon Budget 2022 (Friedlingstein et al., 2022; Hauck et al., 2022). Each GOBM, detailed in Table 1, is driven by meteorological reanalysis and atmospheric  $CO_2$  levels, encapsulating physical, chemical, and biological influences on surface ocean  $pCO_2$  through a system of interconnected differential equations. Each data product is based on multivariate linear regression or machine learning techniques correlating SOCAT observational data with relevant variable observations. The  $fCO_2$  output from each GOBM or data product is provided at a monthly resolution of  $1^\circ \times 1^\circ$ .

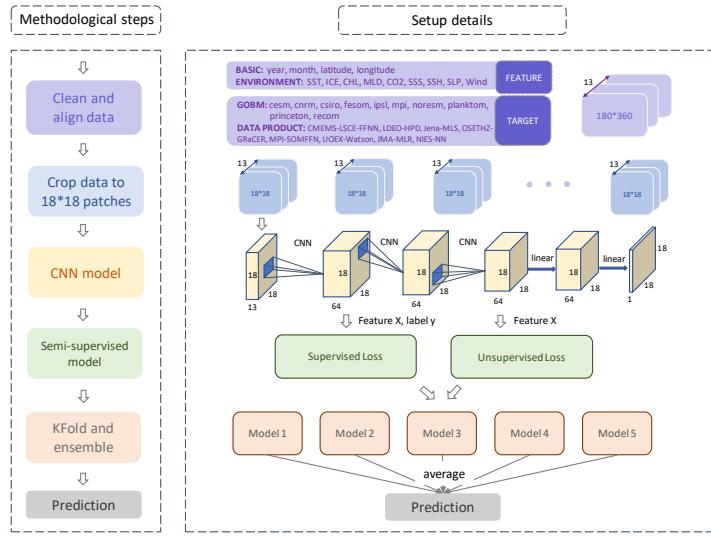


Figure 1: Schematic overview of the methodology and data sources for Carbon Monitor Ocean (CMO-NRT).

### 2.1.2 OBSERVED PREDICTORS

Our predictive variables include biological, chemical, and physical factors linked to  $f\text{CO}_2$  fluctuations. These factors are SST, ICE, SSS,  $x\text{CO}_2$ , MLD, SSH, chl a, SLP, and wind speed. These variables, detailed in Table 2, are bilinearly interpolated from their original grid to a  $1^\circ \times 1^\circ$  monthly resolution to align with our  $f\text{CO}_2$  targets. Given that the  $x\text{CO}_2$  data is only available until the end of 2022, we use a LightGBM model (Ke et al., 2017) to correlate year, month, latitude, longitude, mean atmospheric  $\text{CO}_2$  data and  $x\text{CO}_2$ , enabling near-real-time  $x\text{CO}_2$  data. Data from 1979-2021 is split into training and validation datasets at an 8:2 ratio. Early stopping with LightGBM is implemented, tested on 2022 data, yielding a test RMSE of 1.74, approximately a 0.5% prediction error.

### 2.2 DEEP LEARNING METHOD

Our study devised a deep learning approach tailored for near-real-time estimation of monthly gridded oceanic carbon fluxes, depicted in Figure 1. Integrating inputs of temporal, spatial, and environmental factors, we used GOBM or ocean data product outputs as prediction targets, transforming each dataset into a  $180 \times 360$  grid format. For computational efficiency, all environmental factors were subdivided into  $18 \times 18$  patches.

Recognizing the pivotal influence of surrounding conditions on oceanic carbon absorption, and the proficiency of Convolutional Neural Networks (CNNs) in integrating peripheral data, our model comprises multiple stacked CNN and Linear layers. This allows the capture of both linear and non-linear data relationships. To enhance the model's stability, classic semi-supervised approach, Pseudo-labeling, is employed. For data with labels, the Root Mean Square Error (RMSE) between labels and model predictions was calculated, serving as the supervised loss,  $L_s$ . For unlabeled data points, the RMSE between these pseudo-labels and model predictions was calculated as unsupervised loss  $L_u$ . The final loss of the model is determined by the weighted average of both supervised and unsupervised losses  $wL_u + L_s$ .

To bolster model robustness, we adopted a KFold strategy, partitioning the training data into five subsets. We iteratively used four subsets for training and one for validation, generating five models. Each model yielded a set of predictions for the test data, with the final prediction being the average of all five models' outputs. Detailed model information is provided in the Appendix B.

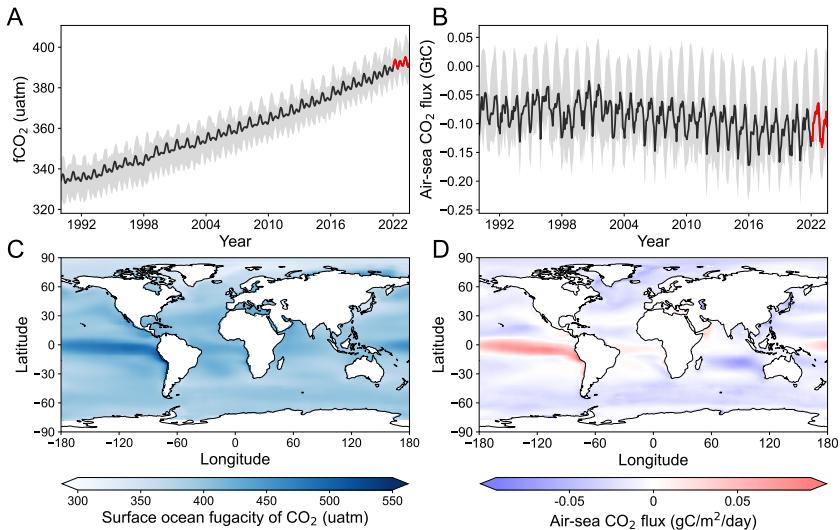


Figure 2: Monthly  $fCO_2$  (A) and air-sea  $CO_2$  flux (positive upward) (B) average of 10 GOBMs and 8 data products over 1990-2021 from Global Carbon Budget 2022 (black lines) and 2022-July 2023 (near-real-time predictions, red lines) from Carbon Monitor Ocean (CMO-NRT). Grey shaded areas represent the range of estimations from 10 GOBMs and 8 data products. Mean  $fCO_2$  (C) and air-sea  $CO_2$  flux (D) over August 2022-July 2023 estimated from CMO-NRT.

Figure 2 displays the monthly oceanic  $fCO_2$  and air-sea  $CO_2$  flux from 1990 through July 2023, alongside their corresponding gridded data for the period of August 2022 to July 2023 as estimated by CMO-NRT. More details of output data can be found at Appendix D

### 2.3 CALCULATIONS OF PCO2 AND AIR-SEA CO2 FLUX

While our deep learning model enables the near-real-time prediction of  $fCO_2$ , a series of further calculations are necessary to determine the amount of  $CO_2$  absorbed by the ocean, resulting in the  $CO_2$  flux. Due to space constraints, the detailed calculation process for this part can be found in Appendix C.

## 3 TECHNICAL VALIDATION

To evaluate CMO-NRT’s near-real-time forecasts, we trained models using data from 2000-2019, excluding the most recent two years’ data from 10 GOBMs and 8 data products, which served as the test set. Our models achieved an RMSE of less than 5% across all targets, with most data products registering around 2% RMSE loss. For more detailed information, please refer to Table 3. We then analyzed the results from three perspectives: correlation, global quantity, and spatial distribution.

- The correlation between CMO-NRT predictions and original outputs for the 10 GOBMs and 8 data products during 2020-2021, was generally strong with most R2 values above 0.9. Scatter plots indicated more stable model performance with GOBMs due to less deviation from the fit line.
- Analysis of global  $fCO_2$  monthly variations revealed a high agreement in seasonality across GOBMs and data products (Figures 4 and 5). Our estimates were slightly higher, but most differences were under 3  $\mu\text{atm}$ . Global aggregate comparisons demonstrated superior performance by GOBMs with most R2 values exceeding 0.85.
- Evaluation of the spatial  $fCO_2$  distribution patterns revealed consistency between CMO-NRT predictions and original outputs across latitudes (Figures 6 and 7), with most discrepancies under 20  $\mu\text{atm}$ . Larger errors were observed in extreme  $fCO_2$  regions, such as the equatorial Pacific, Arctic Ocean, and areas with missing historical data. Monthly

mean fCO<sub>2</sub> values, calculated separately across the 10 GOBMs and 8 data products during 2020-2021, showed high consistency with original results, with generally under 10  $\mu\text{atm}$  discrepancies. Seasonal variations revealed higher original values from GOBMs in the Arctic region during summer, while data product performance was more uniform across months.

## 4 CONCLUSION

Our study, CMO-NRT, introduces a deep learning approach tailored for near-real-time estimation of oceanic carbon sinks, with a minimal delay of 1-2 months. This represents the first global effort for near-real-time oceanic carbon sink estimations. Notably, our method achieves an approximate 2% RMSE prediction for most GOBMs and data products, underscoring its robustness and accuracy. The development of the CMO-NRT dataset effectively bridges the gap between the immediacy of data and the requirement for swift analysis, thereby becoming an invaluable tool for climate change mitigation. Despite significant strides in oceanic carbon sink estimations, the complex evaluation of terrestrial carbon sinks, closely linked with human activities, persists as a lagging issue requiring urgent attention. Drawing on our experience in oceanic explorations, our work could serve as a reference for terrestrial carbon sink predictions, thereby advancing climate change solutions. As the urgency for climate change solutions escalates, our work offers a crucial tool for timely, informed decision-making, playing a central role in global climate change management.

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Table 2: Sources of input data sets.

Variable	Abbreviation	Data Product	Resolution
Sea Surface Temperature	SST	NOAA: OISST Huang et al. (2021)	1981/9-2023/07, daily, 720*1440 (latitude*longitude)
Sea Ice Fraction	ICE		
Sea Surface Salinity	SSS	Met Office: EN4 Good et al. (2013)	1959/01-2023/07, monthly, 173*360
Atmospheric $CO_2$ mixing ratio	$xCO_2$	NOAA: GreenhouseGas Marine Boundary Layer Reference Lan et al. (2023)	1979/01-2022/12, weekly, 180*1
Mixed Layer Depth	MLD	ECMWF: ORAS5	1959/01-2023/08, monthly, 1021*1442
Sea Surface Height	SSH	Copernicus Climate Change Service (2021)	
Chlorophyll-a	Chl a	ESA: GlobColour Maritorena et al. (2010)	1997/09-2023/09, monthly, 180*360
Sea Level Pressure	SLP		
Wind Speed	Wind	ECMWF: ERA5 Hersbach H. et al. (2023)	1959/01-2023/08, monthly, 1021*1442
Year, month, longitude and latitude			

## A DATA

Table 1: Global ocean biogeochemical models and ocean data products used in the Global Carbon Budget 2022 (Friedlingstein et al., 2022; Hauck et al., 2022)

Type	Datasets	Data information
Global ocean biogeochemistry models	NEMO-PlankTOM12 Wright et al. (2021)	Surface ocean fugacity of $CO_2$ , 1959/01-2021/12, monthly, 180*360
	MICOM-HAMOCC (NorESM-OCv1.2) Schwinger et al. (2016)	
	MPIM-HAMOCC6 Lacroix et al. (2021)	
	NEMO3.6-PISCESv2-gas (CNRM) Berthet et al. (2019)	
	FESOM-2.1-REcom2 Hauck et al. (2020)	
	MOM6-COBALT (Princeton) Liao et al. (2020)	
	CESM-ETHZ Doney et al. (2009)	
	NEMO-PISCES (IPSL) Aumont et al. (2015)	
	MRI-ESM2-1 (Nakano et al., 2011; Urakawa et al., 2020)	
	CESM2 Long et al. (2021)	
Ocean data products	MPI-SOMFFN Landschützer et al. (2016)	
	Jena-MLS Rödenbeck et al. (2022)	
	CMEMS-LSCE-FFNNv2 Chau et al. (2022)	
	LDEO-HPD Gloege et al. (2022)	
	UOEx-Watson Watson et al. (2020)	
	NIES-NN Zeng et al. (2014)	
	JMA-MLR Iida et al. (2021)	
	OS-ETHZ-GRaCER Gregor & Gruber (2021)	

## B METHODS

The framework of the model is illustrated in Figure 3. The data points with label values were denoted as  $D_l$ , while those with no labels were termed  $D_u$ . In the initial phase, the model was trained exclusively on  $D_l$  data points. The RMSE between labels and predictions was calculated, serving as the supervised loss,  $L_s$ . Simultaneously, we aimed to ensure stability in predictions even on data without labels. To this end, we employed a classic semi-supervised approach known as pseudo-labeling. Predictions were made by randomly removing 10% of features, which were then used as pseudo-labels. The results predicted by eliminating 30% of features randomly from the input were used as predictions. The RMSE between these pseudo-labels and predictions was calculated as unsupervised loss,  $L_u$ . The model was subsequently updated through backward propagation, utilizing the weighted sum of  $wL_u + L_s$  as the model's loss. This method has been shown to enhance the model's stability.

Our model's detail architecture, as illustrated in the bottom right corner of Figure 3, is an integration of multi-layered CNN and linear models. The model has been meticulously structured to process our input data efficiently and generate accurate predictions.

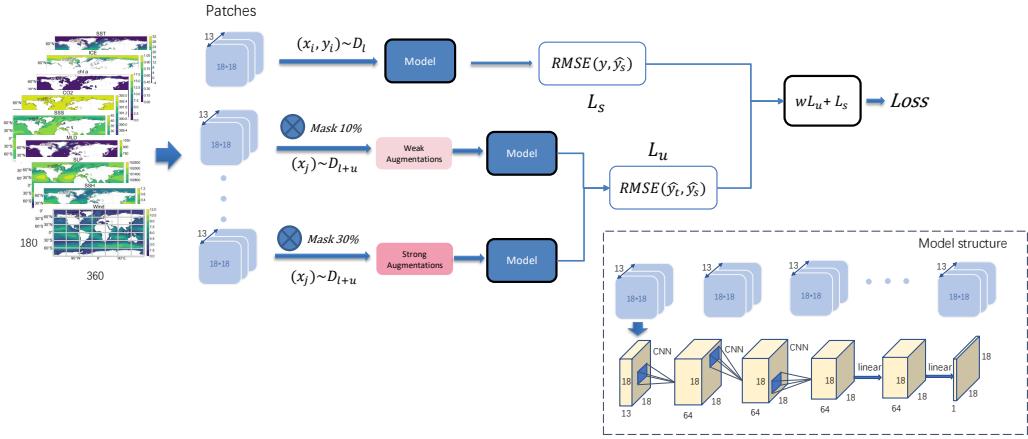


Figure 3: Framework of our methodology.

The input layer is designed to align with the format of our data, preserving a dimension of  $18 \times 18 \times 13$ . This dimensionality accounts for the width, height, and depth (which corresponds to the thirteen input factors considered in this study) and is consistently upheld across all CNN and linear layers. Within the multi-layered CNN models, the hidden layers are configured with dimensions set at 13, 64 and 64. These layers are intentionally designed to automatically and adaptively identify and learn spatial hierarchies from the data. Following the CNN layers, the linear model layers, structured with dimensions of 64, 64, and 1, perform mathematical operations that linearly transform the feature space, aiding in the generation of effective predictions. The architecture of the model culminates in an output layer with a dimension of 1, indicating that it yields a single output value. This value represents our final predicted value for oceanic carbon  $fCO_2$ . The network comprises several layers, including the Convolutional Layer, Rectified Linear Unit (ReLU) Layer, and Fully Connected Layer. These layers work in harmony, each contributing to the processing and transformation of input data, ultimately leading to the output prediction.

## C CALCULATIONS OF $pCO_2$ AND AIR-SEA $CO_2$ FLUX

The  $pCO_2$  is calculated by the following equation:

$$pCO_2 = fCO_2 \times \exp \left( -P_{\text{atm}}^{\text{surf}} \times \frac{(B + 2\delta)}{R \times T} \right)^{-1} \quad (1)$$

where  $P_{\text{atm}}^{\text{surf}}$  is the atmospheric surface pressure from ECMWF Reanalysis version 5 Hersbach H. et al. (2023), T is the sea surface temperature(SST) in Kelvin from National Oceanic and Atmospheric Administration (NOAA) optimally interpolated SST (OISST) Huang et al. (2021), B and  $\delta$  are virial coefficients from Weiss Weiss (1974), R is the gas constant Dickson et al. (2007).

The air-sea  $CO_2$  flux is calculated here by the standard bulk equation:

$$F_{CO_2} = k_w \times S_{CO_2} \times (1 - f_{ice}) \times (pCO_2^{\text{atm-moist}} - pCO_2^{\text{ocean}}) \quad (2)$$

which parameterizes the air-sea  $CO_2$  flux  $F_{CO_2}$  as a function of the gas transfer velocity ( $k_w$ ),  $CO_2$  solubility ( $S_{CO_2}$ ), ice fraction ( $f_{ice}$ ), and partial pressure of  $CO_2$  in moist air ( $pCO_2^{\text{atm-moist}}$ ) and

surface ocean ( $pCO_2^{\text{ocean}}$ ). Solubility is calculated following Weiss Weiss (1974) and partial pressure of moist air ( $pCO_2^{\text{atm-moist}}$ ) is calculated following Equation,

$$pCO_2^{\text{atm-moist}} = xCO_2 \times (P_{\text{atm}} - pH_2O) \quad (3)$$

where  $xCO_2$  is the dry air mixing ratio of atmospheric  $CO_2$  from NOAA Greenhouse Gas Marine Boundary Layer Reference Lan et al. (2023),  $P_{\text{atm}}$  is the total atmospheric pressure from ECMWF Reanalysis version 5 Hersbach H. et al. (2023), and  $pH_2O$  is the saturation vapor pressure Dickson et al. (2007). We use the Wanninkhof Wanninkhof (1992) formulation for the gas transfer velocity:

$$k_w = k_{w,\text{scaled}} \times u^2 \times \left( \frac{S_c}{660} \right)^{-0.5} \quad (4)$$

which parameterizes  $k_w$  as a function of wind speed squared ( $u^2$ ) and the Schmidt number ( $S_c$ ).  $k_w$  is scaled by a factor of  $k_{w,\text{scaled}}$  for each wind product to match the invasion of bomb as of 1994 (Sweeney et al., 2007; Fay & McKinley, 2021). The wind product is from ECMWF Reanalysis version 5 Hersbach H. et al. (2023).

## D RESULT

The CMO-NRT dataset, formatted in Network Common Data Form (NetCDF), provides monthly global oceanic surface  $fCO_2$  and air-sea  $CO_2$  flux data. This dataset spans from January 2022 through July 2023 and is structured into a  $1^\circ \times 1^\circ$  grid. It includes three dimensions and two variables. The dimensions are as follows:

- **Time:** Monthly data, from January 2022 to July 2023.
- **Latitude (lat):** Ranges from  $-90^\circ$  to  $90^\circ$  North.
- **Longitude (lon):** Spans from  $-180^\circ$  to  $180^\circ$  East.
- **Product:** Involves 10 GOBMs and 8 data products.

The two variables covered are:

- $sfCO_2$ : This represents the surface ocean  $fCO_2$ , quantified in units of  $\text{uatm}$ . It is measured across 18 products, 19 time points, and a grid of 180 latitudes by 360 longitudes.
- $fgCO_2$ : This indicates the flux density of total air-sea  $CO_2$  exchange, expressed in  $\text{gC/m}^2/\text{day}$  with a positive value indicating an upward direction. Similar to  $sfCO_2$ , it is measured across the same dimensions. As of the time of this writing, the dataset includes data up to July 2023.

Figure 2 displays the monthly oceanic  $fCO_2$  and air-sea  $CO_2$  flux from 1990 through July 2023, alongside their corresponding gridded data for the period of August 2022 to July 2023 as estimated by CMO-NRT.

## E TECHNICAL VALIDATION

- Overall, the predictions for both GOBMs and data products demonstrated a strong performance, with most having an  $R^2$  value above 0.9. Scatter plot comparisons reveal that, in data products, there are occasional points where predictions significantly deviate from the fit line. In contrast, for GOBMs, most points cluster close to the fit line, indicating more stable performance of the model with GOBMs.
- Global quantity: We further analyzed the monthly variations of global  $fCO_2$ , comparing CMO-NRT predictions with original outputs during 2020-2021 as shown in Figures 4 and 5. The results exhibit a high degree of agreement in terms of seasonality across both GOBMs and data products. Generally, our estimated values were slightly higher than the original values, with most differences being less than  $3 \text{ \mu atm}$ . In terms of global aggregate comparisons, GOBMs demonstrated superior performance, with most  $R^2$  values exceeding 0.85.

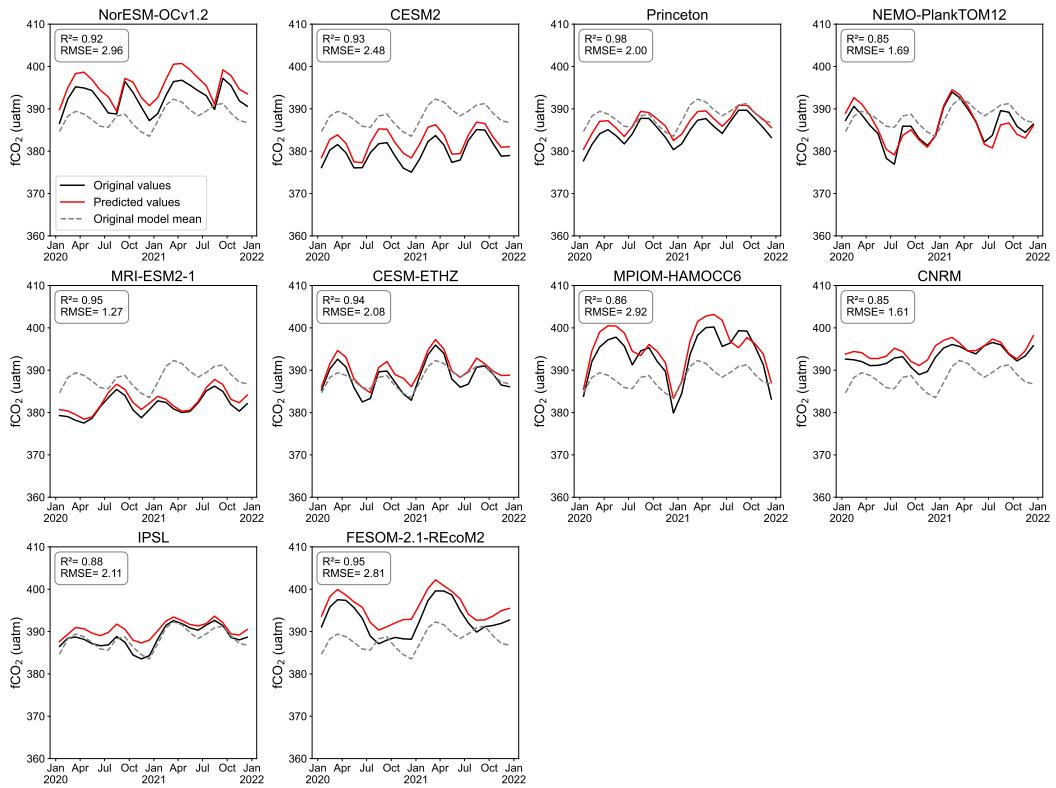


Figure 4: Comparison of global average monthly  $f\text{CO}_2$  between CMO-NRT predictions and original outputs for each of the 10 GOBMs during 2020-2021. The red lines are the predicted results from CMO-NRT. The black lines are the original results from GOBMs. The dashed grey lines are the average results of 10 GOBMs.

- Spatial distribution: We evaluate the spatial patterns of mean  $f\text{CO}_2$  of the CMO-NRT prediction against the original output for each of the 10 GOBMs and 8 data products during 2020-2021 Figures 6 and 7. The spatial variations in  $f\text{CO}_2$  were largely consistent between the CMO-NRT predictions and the original results, with most discrepancies being under 20  $\mu\text{atm}$ . The predicted values also align well with the original outputs in both trend and magnitude across latitudes. Notably larger errors were observed in the equatorial Pacific, including adjacent coastal areas near Peru and Chile to the west and northeast regions near Indonesia and Papua New Guinea, as well as the Arctic Ocean. These errors are primarily due to these regions typically exhibiting extreme maximum and minimum  $f\text{CO}_2$  values. Our model tends to perform less accurately in predicting extremes compared to values closer to the mean, resulting in slightly poorer performance in these areas. Additionally, the MPI-SOMFFN shows lower predictive accuracy in the Arctic region, primarily due to the high frequency of missing historical data in this area.

We also calculated the monthly mean  $f\text{CO}_2$  values by averaging them separately across the 10 GOBMs and 8 data products during 2020-2021. The spatial variation in  $f\text{CO}_2$  showed high consistency between the CMO-NRT predictions and the original results, with discrepancies generally under 10  $\mu\text{atm}$  for GOBMs and data products. The predicted values align well with the original outputs in both trend and magnitude across latitudes. Examining different months, we observed that in the summer months, the original values from GOBMs in the Arctic region were significantly higher than our predictions. In contrast, the performance of data products was more uniform across different months. Typically, the original values from data products were lower than the predictions in the Arctic Ocean and higher in the equatorial Pacific region.

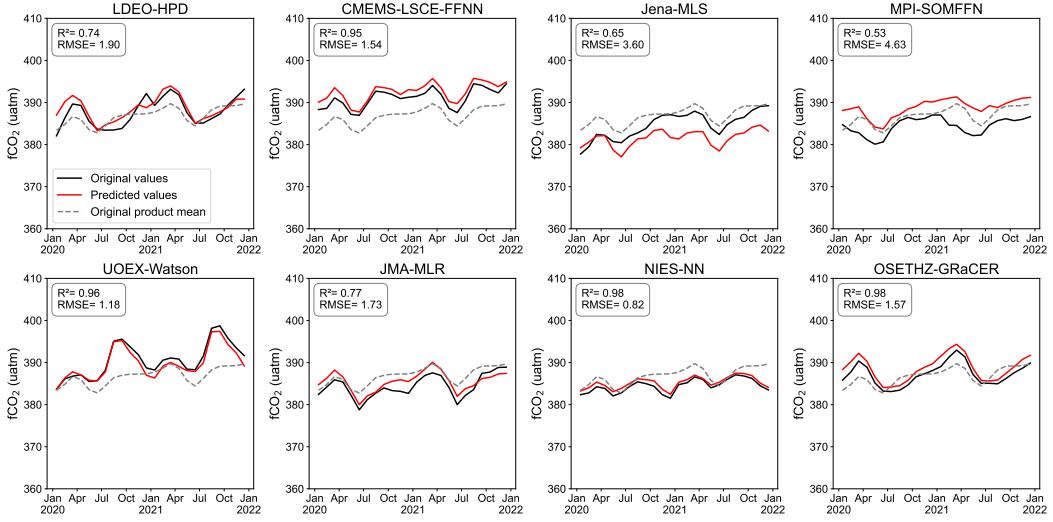


Figure 5: Comparison of global average monthly  $f\text{CO}_2$  between CMO-NRT predictions and original outputs for each of the 8 data products during 2020-2021. The red lines are the predicted results from CMO-NRT. The black lines are the original results from data products. The dashed grey lines are the average results of 8 data products.

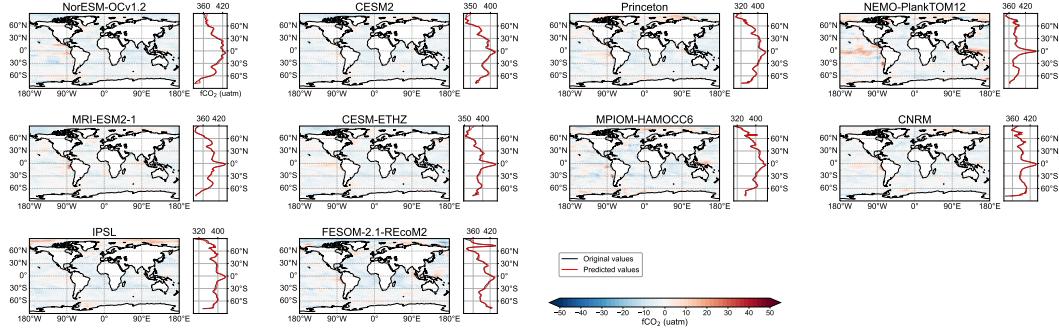


Figure 6: Spatial differences between CMO-NRT predictions and original outputs for each of the 10 GOBMs during 2020-2021. The line graphs represent the values of predictions and original outputs across latitudes.

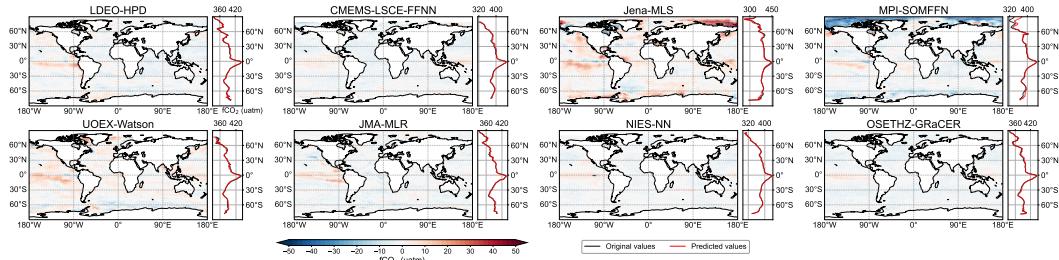


Figure 7: Spatial differences between CMO-NRT predictions and original outputs for each of the 8 data products during 2020-2021. The line graphs represent the values of predictions and original outputs across latitudes.

Table 3: Root Mean Square Error (RMSE) for Different Models  
 GOBM Models

NAME	RMSE	MEAN	NRMSE
NorESM-OCv1.2	10.8	392	2.7%
CESM2	6.8	379	1.7%
MRI-ESM2-1	7.7	381	2.0%
CNRM	11.2	393	2.8%
IPSL	10.6	388	2.7%
MPIOM-HAMOCC6	15.6	393	3.9%
NEMO-PlankTOM12	13.3	386	3.4%
FESOM-2.1-REcoM2	15.1	392	3.8%
CESM-ETHZ	8.7	388	2.2%
Princeton	10.6	384	2.7%

Data Product Models			
NAME	RMSE	MEAN	NRMSE
NIES-NN	5.3	384	1.3%
OSETHZ-GRaCER	6.7	387	1.7%
LDEO-HPD	8.4	387	2.1%
MPI-SOMFFN	16.4	384	4.2%
Jena-MLS	15.1	384	3.9%
UOEX-Watson	16.2	390	4.1%
JMA-MLR	9.9	384	2.5%
CMEMS-LSCE-FFNN	8.4	391	2.1%